

PRELIMINARY GEOLOGIC MAP OF THE WILLITS 7.5' QUADRANGLE

MENDOCINO COUNTY, CALIFORNIA

VERSION 1.0

By

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DESCRIPTION OF MAP UNITS

QUATERNARY SURFICIAL DEPOSITS

af **Artificial fill (historical)**—Consists of engineered and/or non-engineered soil materials; includes larger roadway embankments, earthen dams, levees, and other graded areas.

Qsc **Stream channel deposits (modern to latest Holocene)**—Unconsolidated with active, natural, and constructed stream channels. Composed of loose sand, gravel, and silt. Episodes of bank-fail stream flow are frequent enough to inhibit growth of vegetation.

Qhay **Stream terrace deposits (late Holocene)**—Unconsolidated sand, gravel, silt, and minor clay deposited in active stream settings; typically 5 m or less above the active channel. Commonly inset into adjacent, higher-level deposits of Holocene age.

Qha **Alluvial fan deposits (Holocene)**—Unconsolidated sand, gravel, silt, and minor clay deposited in active or recently active valley fill, inter-fan areas, point bar, and stream settings. Surfaces are relatively flat or gently sloping, minimally dissected except at stream-facing margins, and widely disturbed by agricultural or other human activities. Generally less than 5 m above the active channel.

Qhl **Alluvial fan levee deposits (Holocene)**—Unconsolidated sand, silt, and clay deposited adjacent to a stream channel where water and sediment overtop the channel bank. Mapped based on identification of low, elongated ridges generally parallel to and bordering the channel.

Qhb **Basin deposits (Holocene)**—Unconsolidated alluvium in low-lying floodplain areas in the northern portion of Little Lake Valley, primarily associated with deposition in areas of standing or slowly moving water in topographic basins. Surface is relatively flat or very gently sloping, minimally dissected, and widely disturbed by agricultural or other human activities. Well logs presented by Woolace (2005) indicate primarily clayey sediment with local sand and gravel interbeds. Horizontal stratification is likely. May be interbedded with coarser alluvium from streams that flow into the basin. Interbeds of peat may also be present. Groundwater is shallow in this area of Little Lake Valley and Cardwell (1965) and Woolace (2005) document reports of marshes and a seasonal lake prior to artificial modifications to drainage.

Qhf **Alluvial fan deposits (Holocene)**—Unconsolidated, poorly sorted gravel, sand, and silt deposited chiefly from distributary streamflow and debris flows forming deposits with characteristic fan-shaped morphology that emanate from mountain front drainages and from channels incised through older materials. Fan surfaces show little dissection or soil profile development.

Qdf **Debris fan deposits (Holocene)**—Unconsolidated, poorly sorted gravel, sand, and silt forming relatively steep, fan-shaped deposits at the mouth of small drainages and along steep hillsides where they include undifferentiated colluvium. Sediment derived from debris slides and debris flow events rather than fluvial processes. Limited dissection, poor soil development, and/or absence of mature vegetation indicate relatively recent or ongoing depositional processes. One exposure displays pale-brown to brown silt matrix.

Qa **Alluvial deposits, undivided (Holocene to latest Pleistocene)**—Unconsolidated to weakly cemented sand, gravel, silt, and minor clay, mapped in smaller valleys and where variations in age and/or depositional settings of fans, stream terraces, and active channels are not distinguishable or not delineated at the scale of mapping.

Qf **Alluvial fan deposits, undivided (Holocene to latest Pleistocene)**—Unconsolidated to weakly cemented, poorly sorted gravel, sand, and silt; mapped where morphologic indicators such as incision and soil development suggest the fan surface appears dominant, or the fan age is not distinguishable at map scale.

Qls **Landslide deposits, undivided (Holocene to Pleistocene)**—Unconsolidated surficial failures commonly composed of a mix of silt, colluvium, and weathered bedrock; also includes moderately well-consolidated jumbled rock debris resulting from rock creep and large-scale rotational rockslides. Recognizable by topographic expression or chaotic internal structure. Only larger landlides discernable at the map scale are included. Includes deposits with ages ranging from historically active to older, relict, and heavily dissected slide masses. Queried where landslide existence is questionable.

Qof **Older alluvial fan deposits (early Holocene to late Pleistocene)**—Brown to grayish-brown, slightly to moderately consolidated, poorly sorted gravel, sand, and silt deposited in alluvial fan settings; may also include undifferentiated colluvium from immediately adjacent hillsides. Moderately dissected to deeply incised at confluences with larger, higher-order truncating streams.

Qoa **Older alluvial deposits (early Holocene to late Pleistocene)**—Light yellowish-brown, weakly consolidated, weakly to moderately cemented gravelly sand, silt, and clay deposited in stream and floodplain settings; locally also includes alluvial fan deposits where not mapped separately. Clasts range from rounded to angular locally. Localized gravel exposures are matrix supported and friable. Deposits have been uplifted or otherwise removed from the locus of recent sedimentation preserved in terraces above recently active floodplains and adjacent drainages. Surfaces are dissected to varied degrees, with a moderately developed soil profile. Deposits are dark yellowish-brown with depth, and gravel clasts have clay films. Radiocarbon dates reported by Woolace (2005) suggest deposits exposed in the channel of Haehl Creek are latest Pleistocene; however, that unit would be younger than the Ukiah formation so it is shown as Qoa on the map. Conglomerates on the margins of Little Lake Valley are commonly poorly exposed and the difference between Qoa and Qfu conglomerates can be subtle due to similar lithologies and clast provenance. Unit labels are queried where inaccessible and/or identity is uncertain.

Very old alluvial terrace deposits and surfaces (Pleistocene)—Eroded remnants of fluvial, alluvial fan, and colluvial deposits on isolated strand terrace surfaces cut into the Ukiah formation and bedrock of the Franciscan Complex. Deposits consist of deeply weathered, variably cemented, silty to clayey sand and gravel; original thickness uncertain, now reduced to a thin remnant. The weathering profile locally extends into the strath surface. Gravel content tends to increase down slope. Deposits are generally poorly sorted with variably rounded Franciscan clasts derived from gravelly hillsides. The deposits are divided into subunits inferred from increasing age (1-youngest to 2-oldest) distinguished based on increasing topographic position above adjacent active floodplains, depth of weathering, and soil profile development. Progressive aging of the soil profile is marked by a diminishing to absent organic (A) horizon, development of the argillic (B) horizon with increasing pedogenic clay content, structure, weathering rinds on clasts, and depth and degree of oxidation reddening (Munsell soil color chart colors below classified with dry soil).

Very old alluvial terrace deposits, unit 1 (Pleistocene)—Thin organic horizon over gravel with strong-brown (7.5YR 4/6) weathered sand, silt, and clay matrix. Clasts up to 8 cm in diameter. Yellowish-red (5YR 5/6) clay films on clasts. Relatively youngest subunit representing deposits on the lowest strand terrace.

Very old alluvial terrace deposits, unit 2 (Pleistocene)—Little to no organic horizon; weakly developed argillic horizon over gravel with yellowish-yellow (5YR 6/6) clay matrix. Clasts are commonly 2 to 8 cm in diameter. Reddish-brown (5YR 4/4) clay films on clasts. Moderately indurated.

EARLY QUATERNARY TO LATE TERTIARY CONTINENTAL BASIN DEPOSITS

Ukiah formation (Pleistocene to Pliocene) (?)—Pebble to cobble-bearing conglomerate with interbedded silty sandstone and clayey siltstone. Deposits are well consolidated and generally moderately indurated, with local well-cemented sections. The conglomerate typically appears clast supported and massive or crudely stratified; cross-stratification with coarse channel lag deposits and clast imbrication displays locally. Woolace (2005) interprets a general paleoflow direction to the south-southwest, based on measurements of clast imbrication. Clasts are mostly subrounded to well rounded; Woolace (2005) describes 60 cm boulders in outcrop near southern map boundary. Material appears entirely derived from the Franciscan Complex, dominated by sandstone, with lesser metavolcanic rock, chert, and vein quartz. Fine-grained units are commonly very poorly preserved. Light yellowish-brown oxidation with yellow staining and light-gray mottling locally. Deeper, unoxidized portions described from boreholes and localized exposures exhibit a distinctive blue-gray color. Bedding is generally gently to moderately tilted, except near the Macama fault where it steepens. Woolace (2005) estimates the unit has a minimum stratigraphic thickness of 1100 m in Little Lake Valley based on borehole and stratigraphic data. Woolace (2005) also reports on tephra occurrences identified in the area, including: Thermal Canyon ash (740 ka), Rockland ash (~575 ka), and an unnamed unit correlative to tephras in the Clear Lake basin (~110 ka). Based on stratigraphic data, Woolace (2005) estimates the maximum age of the section is ~1.2 Ma. Cardwell (1965) suggests this unit may be as old as Pliocene; however, the age of the lower boundary of the Pleistocene has since been redefined (Gibbard and others, 2010) and increased approximately one million years; therefore, the basal age may no longer extend into the Pliocene. Ukiah formation is an informal name adopted by CGS for related deposits found in the Ukiah Quadrangle (Deltate and Rubin, 2020), names previously applied to this unit include Ukiah beds (U.S. Army Corps of Engineers, 1955). Continental deposits (Cardwell, 1965), and Capella gravels (Orchard, 1979). Conglomerates on the margins of Little Lake Valley are commonly poorly exposed and the difference between Qoa and Qfu conglomerates can be subtle due to similar lithologies and clast provenance. Unit labels are queried where inaccessible and/or identity is uncertain.

Airport gravels of Woolace (2005) (Pleistocene to Pliocene) (?)—Pebble to cobble-conglomerate, with interbedded laminated siltstone and/or mudstone; well consolidated. The conglomerate typically appears matrix supported and massive or crudely stratified. Clasts are generally subrounded to well rounded, and range in size up to 25 cm. Material appears derived from the Franciscan Complex; dominated by sandstone, with lesser metamorphic rock and chert, and includes distinctive white mica quartzite (?) of uncertain origin. The unit is generally deeply weathered with yellow to brownish-yellow colors; sandstone clasts and other fine-grained clasts are reduced to a friable condition. Bedding is gently dipping. Unit defined by Woolace (2005) for deposits exposed near Willits Municipal Airport. May be underlain locally by a very old to yellowish-red (5YR 5/6) clay matrix. Age is inferred to be equivalent to Qfu based on the unit having some common characteristics, including the following: similar conglomerate properties; interbedded fine-grained facies; and possible overlying very old terrace deposits. However, the extensive weathering and highly disturbed it from Qfu observed along margins of Little Lake Valley. Woolace (2005) suggests this unit may be older than Pleistocene; however, the age of the lower boundary of the Pleistocene has since been redefined (Gibbard and others, 2010) and increased approximately one million years; therefore, the basal age of the Airport gravels may not extend beyond the Pleistocene.

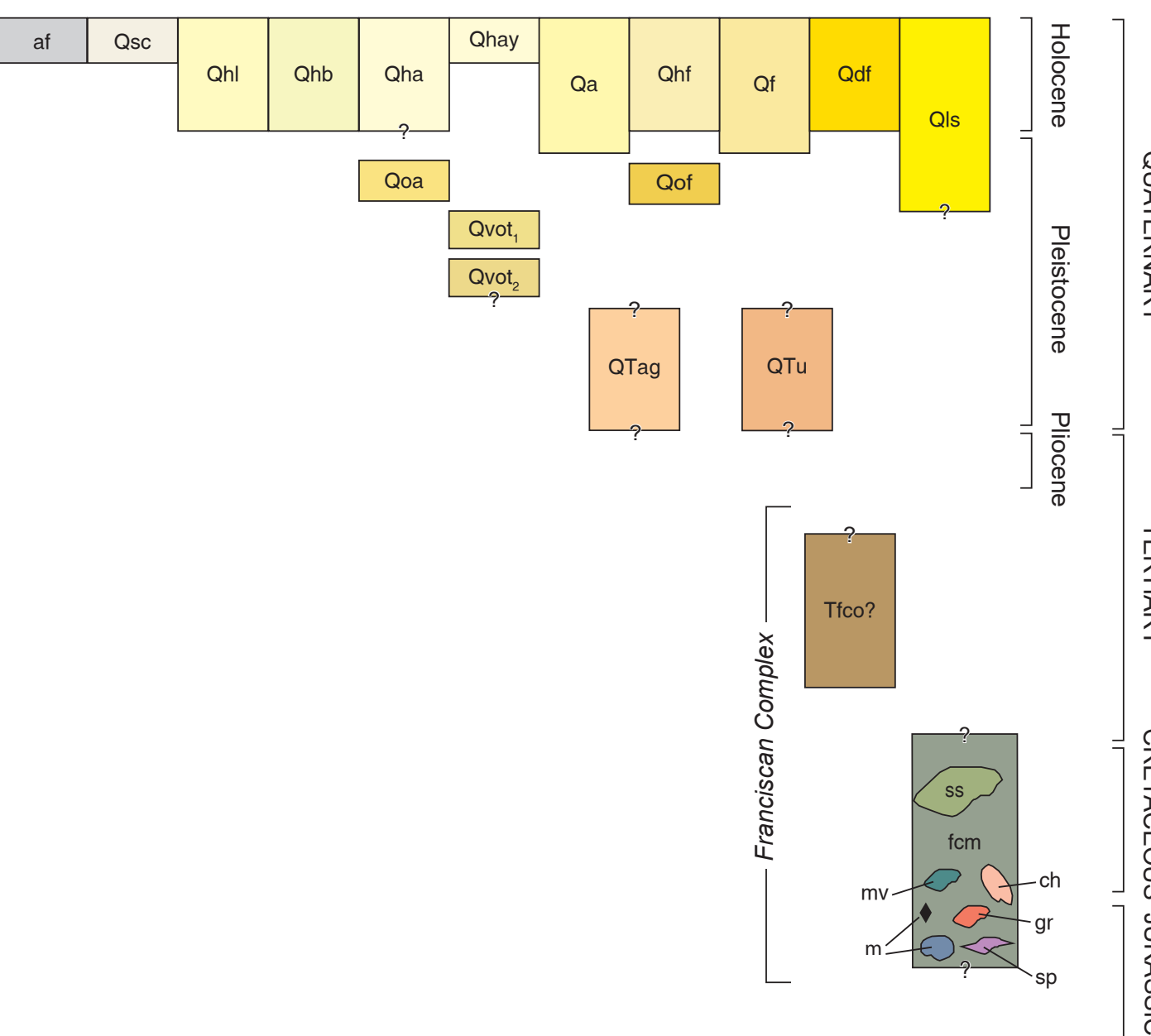
WILLITS 7.5-MINUTE QUADRANGLE

SOURCES OF MAP DATA

1. Durham, 1979*
2. Smith, 1981a
3. Simpson, 2016
4. Cardwell, 1965
5. Farrar, 1986
6. Woolace, 2005
7. Rubin, 2024*

*Data source covers entire quadrangle.

CORRELATION OF MAP UNITS



MAP SYMBOLS

--- Contact between map units. Solid where accurately located; long dash where approximately located; short dash where inferred; dotted where concealed.

--- Fault—line where accurately located; long dash where approximately located; short dash where inferred; dotted where concealed; queried where uncertain. Arrow and number indicate direction and angle of dip of fault plane. Relative horizontal movement shown by arrows parallel to fault.

Landslide—Arrows indicate principal direction of movement. Where mapped as a landslide complex, adjacent defined slides have different relative ages and/or failure types.

Mélange block—See description of map unit "m" for details.

Strike and dip of geologic structure; number indicates dip angle in degrees.

Vertical bedding

Metamorphic foliation

FRANCISCAN COMPLEX – COASTAL BELT

Central Belt, undivided (Oligocene to Eocene)—Very pale-orange, laminated argillite. Includes grayish-orange, poorly sorted, fine to medium-grained arkosic sandstone interbeds. Sandstone beds are 15 cm to >30 cm thick; discontinuous; possibly sheared or deformed. Sandstone includes detrital biotite and mica. Weathered with yellowish-brown oxide stains. Bailey and others (1964) document widespread presence of distinctive detrital potassium feldspar in Coastal Belt sandstones. One occurrence of Coastal Belt material is queried along the west-central margin of the Willits Quadrangle. The identity is queried due to limited exposure and uncertain structural relationships with the adjacent Central Belt. Kilbourne (1984) depicts the Coastal Belt as thrust over the Central Belt in the adjacent quadrangle, in contrast to previous work by Kratzer (1970) showing the opposite relationship. Regionally, the Coastal Belt is understood to have been thrust under the Central Belt (e.g., Kratzer, 1976; Langenheim and others, 2013), resulting in extensive deformation and foliation metamorphism (e.g., Rubin and McLaughlin, 2012). Detrital zircons from sandstones within the Coastal Belt near the study area suggest maximum depositional ages that range from early Oligocene to Eocene (Dumitru and others, 2015).

FRANCISCAN COMPLEX – CENTRAL BELT

Central Belt mélange, undivided (Paleogene (?) Cretaceous and Jurassic) (?)—Pervasively sheared, dark-gray to black, scaly argillite and sandstone (usually referred to as graywacke), which form a matrix enclosing small pods, blocks, and large slabs of more intact rock of various lithologies; overprinted by low-grade metamorphism. The Central Belt matrix sandstone is gray to olive gray, yellowish brown where oxidized, well indurated, poorly sorted, and arkosic; commonly consisting of approximately equal proportions of quartz, plagioclase feldspar, and lithic fragments, with 5 to 20 percent of the rock being matrix (Hock, 1970; Stanford, 1991). Coarse-shale chips and detrital biotite are common lithic components in hand sample. The sandstone lacks potassium feldspar and displays textures and metamorphic mineral assemblages consistent with Central Belt units elsewhere (Joyko and others, 1989; Stanford, 1991). Sandstone appears weakly recrystallized due to metamorphism and is commonly cut by veins of detrital sandstone and shale resembling turbidites; is exposed locally. Due to pervasive deformation and limited exposure, bedding is variable over short distances and is not traceable across a mappable area. The mélange undergoes rolling, characteristically hummocky topography, particularly where isolated blocks within shale matrix are the dominant structure. Hummocky topography is also commonly associated with overprinting of the structure by extensive Quaternary landslide movement. The unit also includes large areas with landslide-like features not distinguishable as individual landlides. Mélanges are associated with originating from tectonic and/or depositional mixing processes, which likely also included coeval subhorizontal transverse movement. Mélange is generally poorly interpreted as faulted, but many are of uncertain origin based on limited available exposures. In the Redwood Valley Quadrangle (Rubin, 2021), chert blocks within the mélange range from Jurassic through Cretaceous in age (Stanford, 1991); however, dates from mélange matrix collected elsewhere in the region suggest tectonic and/or depositional assemblage of the mélange may have extended into the early Tertiary (McLaughlin and others, 2018). Detrital zircons from Central Belt sandstones in the map region indicate an approximate mid-Cretaceous maximum depositional age, with ages ranging from 107 to 85 Ma (Dumitru and others, 2015). Other deposits within the undifferentiated mélange may extend into the Jurassic. Portions of undifferentiated mélange in this study may include additional distinctive mappable blocks and slabs of lithologies described below, but wide areas were largely inaccessible in the field. Mappable units within the mélange include:

Sandstone—Map unit combines isolated blocks and areally extensive occurrences of poorly to moderately sorted, fine to medium-grained, low-grade metamorphic arkosic sandstone; includes rare conglomerate. Commonly light olive gray in color; ranges from diskly yellow, to grayish olive, to greenish gray and medium gray where minimally weathered. Lithic component varies in abundance; detrital biotite and coarse-shale chips are common. Bedding varies from massive to interbedded sandstone and shale. Rock exposures vary widely from hard and cemented, with little oxide staining, to intensely weathered and reduced to a friable, grayish-orange and reddish-yellow, highly oxidized condition in exposures along Ridgewood Road towards the southeast corner of the quadrangle. Degree of internal deformation varies from relatively coarse zones of shearing or other deformation, to broken. May locally include blocks of other subunits. Generally weakly recrystallized, consistent with metamorphic Textural Zone 1 of Blake and others (1967), higher grade, foliated metamorphic units found locally. Mapped where the unit is large enough to meet minimum map unit size, and where assignment to another map unit is uncertain. Relative to Central Belt mélange (fcm) matrix, larger mapped occurrences of this unit are generally more coherent, contain fewer exotic lithologies, are less prone to large landlides, and are underlain by steeper more resistant slopes, as expressed in the northwest corner of the quadrangle. Generally associated with denser forest vegetation relative to fcm.

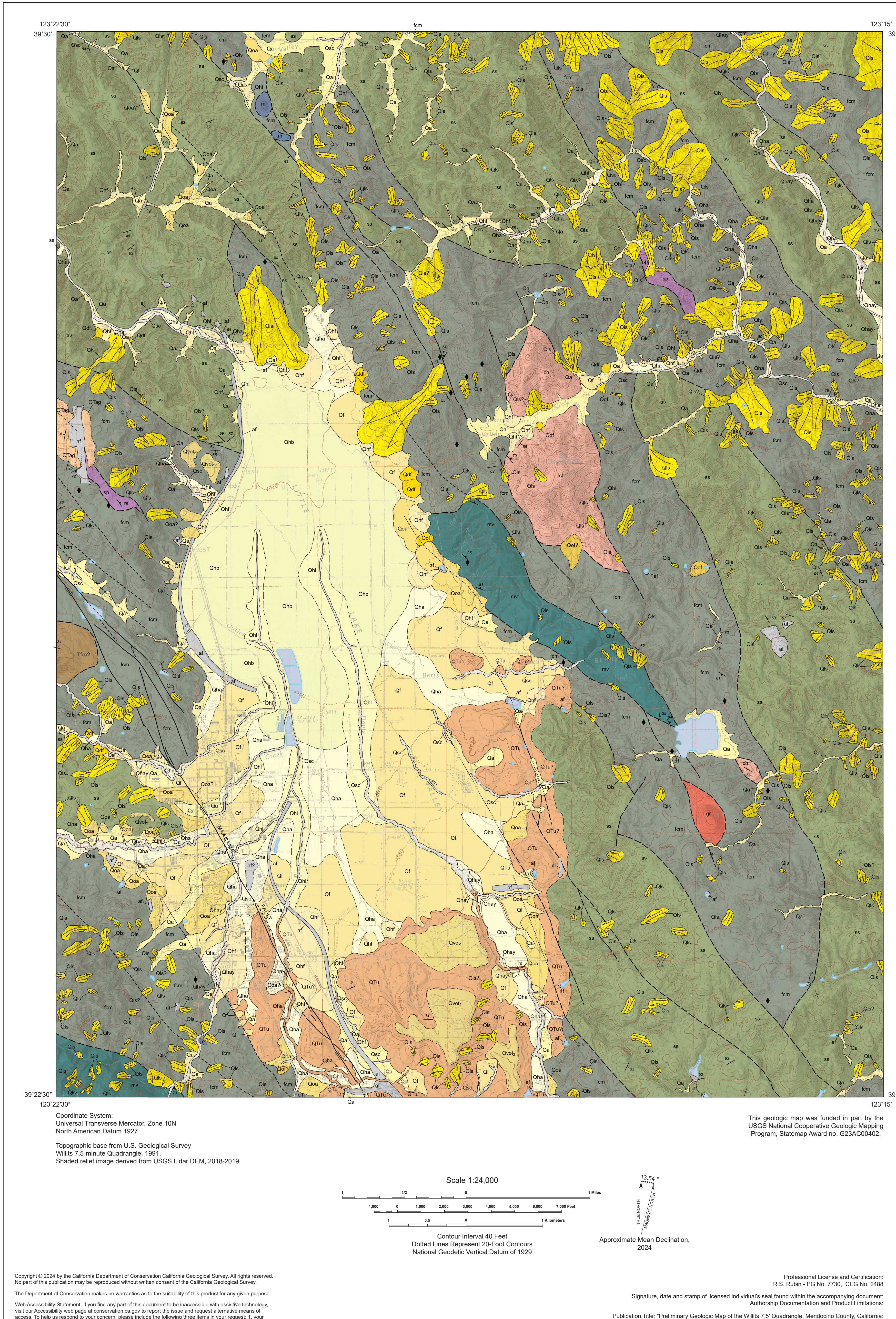
Chert—Red, green, and white chert; variably metamorphosed. Occurs mostly in rhythmically bedded packets with dark-gray shale partings and locally contorted bedding; includes some sandstone. Locally massive or as brecciated fragments; also present as isolated mélange blocks too small to distinguish at map scale. Large section delineated near center of map is associated with resistant topography and steep slopes. Chert samples from various units in the Redwood Valley Quadrangle (Rubin, 2021) yielded radiolarians ranging from Tortonian (late Early Jurassic) to Albian (late Early Cretaceous) in age (Stanford, 1991).

Metavolcanic rock (greenstone)—Olive gray to greenish-black, strong-brown and yellowish-red weathering, aphanitic, variably weathered and altered, mafic volcanic rocks; predominantly massive and bedded; variably sheared; locally stained black. Section mapped above east side of Little Lake Valley is intensely deformed and fractured; mapped areas include some local sandstone and other lithologies. Also present as isolated mélange blocks too small to distinguish at map scale. Weathers to a soil with distinctive red and brown colors.

High-grade metamorphic rock—Metamorphic rock with phyllitic to gneissic texture; largely blueschist grade although other tectonically metamorphosed Franciscan rock types are included; mapped as polygons where unit exposures are large enough to meet minimum map unit size; black diamond symbols are used for discrete mélange blocks smaller than minimum map unit size.

Serpentine—Pale-olive to grayish-olive-green, serpentinized ultramafic rocks. Commonly pervasively sheared and foliated; ultramafic protolith texture is preserved locally. Appears to enclose blocks of other fine lithologies locally. Weathers to regolith and soil with distinctive blue-green or reddish colors. Occurs as mappable units and isolated blocks within the mélange, also along shear zones. Vegetation on serpentine is chiefly sparse shrubs.

Granite—Greenish-gray, medium-grained felsic gneissoid. Weakly recrystallized and foliated. This unit is interpreted to be a mélange block. Includes possible blue amphiboles, indicative of higher-grade metamorphism than the surrounding mélange matrix. Unit is associated with distinctive smooth, steep slopes and an overall domed expression. Plutonic mélange blocks in the Franciscan are reported to the south by Erickson (2011), with one having a reported U/Pb zircon age of 165 ± 1 Ma.

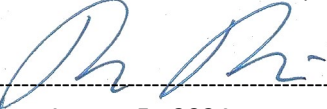


AUTHORSHIP DOCUMENTATION AND PRODUCT LIMITATIONS

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