

# Coordinate System:

#### Universal Transverse Mercator, Zone 11N, North American Datum 1927 Topographic base from U.S. Geological Survey, Crescent Peak 7.5-minute quadrangle, 1984.

Shaded relief image derived from USGS Lidar DEM, 2019.

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# Scale 1:24,000 1,000 0 1,000 2,000 3,000 4,000 5,000

Contour Interval: 10 Meters National Geodetic Vertical Datum of 1929



	115° 7′ 30″	
3750 V VUE 2-24 13 13750 V VUE 2-24 13 Diversion in barrier and the second se	35° 30'	PRELIMINARY GEOL NEW YORK MC
TISET Prospect Contract Contra		
		Qc       Colluvial deposits (Holocene)—Unconsolidated to weakly consolidated slope deposits consisting of colluvial debris and talus derived from up-slope outcrops. Some deposits are degraded by erosion and likely older Holocene in age. Distinguished only where deposits are prominent or conceal relationships between underlying units.         Qw       Modern Wash deposits (late Holocene)—Unconsolidated sand, gravel, cobbles, and local boulders deposited in recently active stream channels. Occur as narrow deposits in canyons upstream of the mountain front and as anastomosing to elongate deposits where active flow paths continue beyond the mountain front and traverse older fan deposits, but do not form fan-shaped landforms; locally includes small areas of Qyf deposits. Sediments are generally derived from local bedrock or reworked from adjacent older Quaternary deposits. Materials subject to mobilization an redenosition during storm events and therefore lack oxidation on clasts and only support local sparse vegetation.
Ball Ball Ball Ball Ball Ball Ball Ball	34 5 A S A S A S A S A S A S A S A S A S A	Qf Modern alluvial fan deposits (late Holocene)—Unconsolidated to weakly consolidated, poorly sorted, sand, gravel, cobble, and local boulder deposits with intermixed silt forming active, essentially undissected, alluvial fans. Fan apice commonly occur at or upstream of the mountain front, and locally form outboard of the mountain front and downgradie of incised older fans. Sediments subject to mobilization and redeposition during storm events and therefore only support local, sparse vegetation. Clasts are derived from up-slope rock sources and reworked from adjacent older fan deposits; clasts typically unweathered with little to no oxidation or desert varnish. Fans are typically braided and include a composite mix of sediment-rich stream deposits and poorly bedded and poorly sorted debris flow deposits, containing angular to sub-angular pebble- to boulder-size clasts closer to the mountain front.
	14537 295	Qfv       Modern alluvial fan deposits, volcanic clast dominated (late Holocene)         Qa       Modern alluvium (late Holocene)—Unconsolidated to weakly consolidated, mostly undissected, sand, silt, and pebb
		to cobble gravel, recently deposited parallel to localized stream valleys and/or spread more regionally onto broad alluv valleys.
A Crippled Jack the A Crip		QywYounger wash deposits (middle Holocene to late Pleistocene)—Weakly consolidated sand, gravel, cobbles, and local boulder deposits. Commonly occurs as elongate deposits on the margins of active Qw deposits in larger canyons upstream of the mountain front. Sediments are generally derived from bedrock in upstream source areas. Sediments m be subject to mobilization and redeposition during large storm events. Vegetation may be sparse to moderately dense.QyfYounger alluvial fan deposits, undifferentiated (middle Holocene to late Pleistocene)—Unconsolidated to weakly
		consolidated, poorly sorted, sand, gravel, cobble, and local boulder deposits with intermixed silt, generally lacking pedogenic carbonate, form undissected to slightly dissected alluvial fans with ubiquitous small-scale surface roughness. Cones and broad aprons with fan apices typically occur upstream of the mountain front in the larger canyons but also occur at or outboard of the mountain front downstream of older incised fans. Locally includes narrow active Qf/Qw deposits and small areas of Qof. Sediments may be subject to local mobilization and redeposition during large storm events; vegetation may be sparse to moderately dense. Clasts are derived from up-slope sources and reworked from adjacent older fan deposits; clasts are typically weakly weathered with weak oxidation and desert varnish. Fans are composed of a composite mix of sediment-laden stream deposits and poorly bedded, poorly sorted debris flow deposit containing pebble- to boulder-size clasts closer to the mountain front
		Qyfv       Younger alluvial fan deposits, volcanic clast dominated (middle Holocene to late Pleistocene)
		Qyf2Younger alluvial fan deposits, younger facies (middle Holocene to late Pleistocene)Younger facies of Qyf the same undissected to weakly dissected with sparse vegetation and weak oxidation of clasts. Sediments of Qyf2 are subject to mobilization and redeposition during large storm events and flow avulsions.Qyfv2Younger alluvial fan deposits, younger facies, volcanic clast dominated (middle Holocene to late
Que Xmg Xig? Qot Xig? Que Cof av Xig?		Qyf1       Pleistocene)         Qyf1       Younger alluvial fan deposits, older facies (middle Holocene to late Pleistocene)—Older facies of Qyf that is slightly to moderately dissected with degraded/smoothed small-scale roughness compared to Qyf2, sparse to moderate vegetation density, and weak to moderate oxidation of clasts. Deposits mapped as Qyf1 likely represent more than one age of deposition and may overlap in age with deposits mapped as Qof, particularly on the east flank of the New York Mountains. Qyf1 locally overlies Qof or Qvof, particularly in the Pinto Valley Quadrangle.
Quy Qof		Qyfv1       Sediments in this unit are rarely subject to mobilization and redeposition by storm runoff.         Younger alluvial fan deposits, older facies, older facies, volcanic clast dominated (middle Holocene to la Pleistocene)
Qof Qof Qof Qof Qof		Qya Younger alluvium (Holocene and late Pleistocene)—Unconsolidated silty sand, grusy sand, gravel, cobbles, and loc boulders. Deposited in canyon areas upstream of the mountain front, either in broad alluvial valleys or in narrower valleys where active wash deposits are subordinate or lacking. Alluvial surfaces are slightly elevated from active channels and support common vegetation, suggesting they are only inundated during larger, uncommon flood events. Clasts are generally unweathered with little oxidation. Locally more than one level of younger alluvium may be preserved.
Qof Qyf2 Qw Qof Xlg Qyf2 Qw Qof Xlg Qyf2 Qw Xlg Qvof. Qvof. Xlg Qvof. Xlg Qvof. Xlg Qvof. Xlg Qvof. Xlg Qvof. Xlg Qvof.		Qof Older fan deposits (late to middle Pleistocene)—Slightly to moderately consolidated, poorly sorted, silty, pebbly sat to coarse gravel and boulder fan deposits. Broad to isolated fan surfaces are typically smooth to moderately dissected a isolated by intervening younger fan deposits; surfaces support stable vegetation and clasts exhibit moderate oxidation/ desert varnish patinas and desert pavement. Surfaces are generally elevated at least several meters above active channed grade, and not subject to historic flood inundation. The underlying bedrock is locally exposed around margins of denosits. One fine comparise the print Valley Ouedrangle and exposed due to generally younger.
Qof XIp Xig Qof Qof Qof Qof + Xfp? Qof XIp Xig Qof Qof Qof Qvof + Xfp? Xip 60 Xbt? Xig Qvfz Xig Xfp? Qvfz Qvfz Qvfz Qvfz Qvfz		incision. Deposits occur both upstream and downstream of the mountain front and typically range from less than 1 m to about 5 m thick on the west flank of the New York Mountains; on the east flank of the range deposits tend to be thicker; deposits mapped as Qoa in the Watson Wash drainage in the Pinto Valley Quadrangle range up to 40 m thick and may be a distinct older unit that includes old wash deposits. Deposits include light to moderate pedogenic carbona development along selected beds and fractures, which is typically weakly indurated. An IRSL date of 70.09 +/- 7.27 ka was determined for sample CRBS008 located on west flank of New York Mountains in the Crescent Peek Quadrangle
XIP Qyfz Job Qof 78/Xbl? Qof Qyw	2012	Qofv       Older fan deposits, volcanic clast dominated (late to middle Pleistocene)
Xtp Qyts Xlg 70.1+/-7.3ka Xlg Qyts Xlg Xtp Xtp Qyts Xtg	Xa Xlg Xa Xlg	Qvof Very old fan deposits (early Pleistocene)—Moderately to well-consolidated silt, sand, gravel, cobbles, and boulders locally exceeding 1 m in diameter and forming deposits up to 15 m thick; some clasts are exotic to local source areas. Fan surfaces are smooth or rounded and degraded such that the original surface and surface oxidation patinas are commonly lost; eroded surfaces range in color from light gray to light brown. Fan surfaces may extend up to tens of meters above adjacent channel grade and underlying bedrock may be exposed. Deposits generally found a short distant
Xigr Xigr 42 Xigr Xig 62 Qyti Qyti Qyti	Xgnt Qw Xgnt Xgnt Xgnt Xgnt 35° 23' 49"	QTg Ancient gravel deposits (Pleistocene to Pliocene)—Light-gray weathering, moderately indurated, where observed. and sand deposits with locally extensive pedogenic carbonate. Deposits are highly dissected and degraded such that the original alluvial surfaces and oxidation patinas are commonly lost. Deposits range up to 20 m thick and are less related
fp xlg xlg Qw Xa Qw Xmg Qof? to to to to to to to to to to	80 Xmg Xgnt Xgnt Xmg	Quadrangle, both on the southeast flank of the New York Mountains and where they locally extend westward across the range crest and infill an ancient drainage. The deposit is also extensive on the northern portion of the Crescent Pea Quadrangle where they overlie Miocene volcanic rock and are commonly tilted down to the southeast and locally cut I faulting.
Xa Xmg Qof Qof	Xmg Qw	TERTIARY UNITS
Qof Qof Xmg Xmg Xmg Xmg Xmg Xmg Xmg	Qc Xlg Xlg Xlg Xlg Xlg Xlg Qya? Xgnt Tab Qya?	19 Gravels (Whotehe)—Whotehely to well consolidated huvial boulder to pebble gravel to coarse debris-flow deposits; local avalanche breccia and gravity-slide breccia sequences. Exotic clasts of porphyritic and augen gneiss with boulde up to 3 m in diameter. Sand interbeds of siliciclastic, volcaniclastic, and locally arkosic makeup. Caliche development and clast types not derived from local sources are characteristic of unit, indicating through-going fluvial environments and (or) far-travelled breccia sheets. Notable lack of volcanic clasts. Locally interfingers with Tab in the Crescent Peal quadrangle area.
Alg     Xmg     Xmg     Qyw       Xlg     Xmg     Xmg     Xmg       Xlg     Xmg?     Xlg       Xfp?     Xmg?     Xlg       Xmg?     Td     Tab	Tab	TabAndesite and basalt (Miocene)—Andesitic block and ash flow tuffs, volcanic breccias, basalts, and andesites. Block and ash flow tuffs are composed of $40-80\%$ , cm- to m-sized fragments of basaltic andesite and hornblende plagioclase andesite. Hornblende and plagioclase may be present in the matrix, and each may comprise up to 10% of unit. Volcanic breccias include lahar deposits, and contain $40-60\%$ , cm- to 0.5 m-sized blocks of mixed volcanic rock fragments (basalt, hornblende andesite) and lesser intrusive lithic fragments, in matrices of mixed ash, mud, and sand. Large, roughly 10 m outcrops of andesite and basalt may display features such as stretched vesicles, autobrecciation, and columnar joints. Secondary mineralization is sometimes present and includes quartz amygdules and carbonate cementation of breccia matrices. Samples yield ages of $17.6 \pm 0.4$ Ma (K, Amplexicater) contract of Vinter, 19.9 mineralization.
Contraction of the set	35° 22' 30" 115° 7' 30"	<ul> <li>Commentation of orecera matrices. Samples yield ages of 17.6 +/- 0.4 Ma (K-Ar plagioclase) southeast of Nipton, 18.8 + 0.5 Ma (K-Ar biotite) and 15.6 +/- 0.4 Ma (K-Ar plagioclase) southeast of Crescent Peak (Miller and Wooden, 1993).</li> <li>Dacite (Miocene)—Andesitic block and ash flow tuffs and lahar breccias. The block and ash flow tuff contains 1cm</li> </ul>
This geologic map is based upon work su under Cooperative Agreement No. G19A0	upported by the U.S. Geological Survey C00304.	to 1m blocks of hornblende andesite in an ash matrix and includes thin layers of welded tuff. Hornblende is present in the blocks and matrix and may comprise up to 15% of unit. Lahar breccia contains 50%, cm to m sized volcanic rock fragments (including hornblende andesite, basalt, and intermediate to mafic anhanitic fragments that are altered to pur

The views and conclusions contained in this document are those of the authors and should not be interpreted as representing the opinions or policies of the U.S. Geological Survey. Mention of trade names or commercial products does not constitute their endorsement by the U.S. Geological Survey.

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1990).

metamorphism.

PROTEROZOIC BASEMENT UNITS Diabase (Mesoproterozoic)—Dark-green to brown, altered, subophitic diabase. Occurs as dikes in map area. Composed of 50–60% plagioclase and 40–50% clinopyroxene. Dated at 1100 Ma (Howard, 1991). Granodiorite of Crippled Jack Well (Paleoproterozoic)—Dark-gray to brown, hornblende-biotite granodiorite with bundant mafic minerals (up to 40%). Accessory garnet locally. Dominantly porphyritic, locally mylonitic. Salt-andpepper appearance on fresh surfaces, tan to brown on weathered surfaces. Miller and Wooden (1994) report a U-Pb date of 1659 +/-4 Ma on zircon collected at Moore siding and a second U-Pb date of 1,662 +/-7 Ma was obtained near the type locality of the pluton (Miller pers com, Dec. 2021), both based on TIMS analyses. New U-Pb dating of rock mapped as Xcj just north of the Nevada border produced preliminary ages of greater than 1,700 Ma. Analyses were conducted on both titanite and zircon grains mounted in thin section using laser ablation ICPMS analyses at the CSUN Laser Lab. Additional U-Pb analyses are in progress to further assess the age of Xcj crystallization and subsequent deformation and

# IMINARY GEOLOGIC MAP OF THE CRESCENT PEAK 7.5' QUADRANGLE, NEW YORK MOUNTAINS, SAN BERNARDINO COUNTY, CALIFORNIA

VERSION 1.1

David Reioux, Benjamin Parrish, Brian J. Swanson, and Howard J. Brown<sup>†</sup>

GIS and Digital Preparation by Deshawn A. Brown Jr., Jeremy L. Altringer, and Milton Fonseca

2022

## OF MAP UNITS

#### **ICIAL DEPOSITS**

its are degraded by erosion and likely older Holocene in age. nceal relationships between underlying units. lated sand, gravel, cobbles, and local boulders deposited posits in canyons upstream of the mountain front and as aths continue beyond the mountain front and traverse older fan y includes small areas of Qyf deposits. Sediments are generally older Quaternary deposits. Materials subject to mobilization and idation on clasts and only support local, sparse vegetation. nsolidated to weakly consolidated, poorly sorted, sand, gravel, forming active, essentially undissected, alluvial fans. Fan apices and locally form outboard of the mountain front and downgradient n and redeposition during storm events and therefore only support ope rock sources and reworked from adjacent older fan deposits; or desert varnish. Fans are typically braided and include a oorly bedded and poorly sorted debris flow deposits, containing

#### inated (late Holocene)

weakly consolidated, mostly undissected, sand, silt, and pebble l stream valleys and/or spread more regionally onto broad alluvial

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Jnconsolidated silty sand, grusy sand, gravel, cobbles, and local ountain front, either in broad alluvial valleys or in narrower acking. Alluvial surfaces are slightly elevated from active they are only inundated during larger, uncommon flood events. Locally more than one level of younger alluvium may be present. htly to moderately consolidated, poorly sorted, silty, pebbly sand lated fan surfaces are typically smooth to moderately dissected and upport stable vegetation and clasts exhibit moderate oxidation/ generally elevated at least several meters above active channel underlying bedrock is locally exposed around margins of Valley Quadrangle and exposed due to geologically young of the mountain front and typically range from less than 1 Mountains; on the east flank of the range deposits tend to be ainage in the Pinto Valley Quadrangle range up to 40 m thick deposits. Deposits include light to moderate pedogenic carbonate typically weakly indurated. An IRSL date of 70.09 +/- 7.27 ka lank of New York Mountains in the Crescent Peak Quadrangle.

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y to well-consolidated silt, sand, gravel, cobbles, and boulders s up to 15 m thick; some clasts are exotic to local source areas. that the original surface and surface oxidation patinas are ght gray to light brown. Fan surfaces may extend up to tens of edrock may be exposed. Deposits generally found a short distance

#### UNITS

ash flow tuffs, volcanic breccias, basalts, and andesites. n- to m-sized fragments of basaltic andesite and hornblende be present in the matrix, and each may comprise up to 10% of tain 40-60%, cm- to 0.5 m-sized blocks of mixed volcanic rock usive lithic fragments, in matrices of mixed ash, mud, and sand. y display features such as stretched vesicles, autobrecciation, imes present and includes quartz amygdules and carbonate f 17.6 +/- 0.4 Ma (K-Ar plagioclase) southeast of Nipton, 18.8 +/ clase) southeast of Crescent Peak (Miller and Wooden, 1993). and lahar breccias. The block and ash flow tuff contains 1cm nd includes thin layers of welded tuff. Hornblende is present in unit. Lahar breccia contains 50%, cm to m sized volcanic rock termediate to matic aphanitic fragments that are altered to purple and red) in a matrix of ash and sub-rounded, medium- to coarse-grained sand.

Peach Spring Tuff (Miocene)—Pink to tan vitric rhyolitic tuffs and volcaniclastic sandstones. Pink tuffs are welded and contain variable amounts of fiamme with lesser abundant (~3%) volcanic lithic fragments and crystals (biotite, hornblende, and abundant sanidine). East of Ivanpah tuffs may be unwelded and tan in color. Sandstone underlies tuffs and is composed of medium- to coarse-grained quartz sand, with lesser abundant ash and dark lithic fragments, including volcanic rock fragments and fiamme. Ar/Ar dating of sanidine yield an age of 18.5 +/- 0.2 Ma (Nielson and Others,

Leucocratic granite (Paleoproterozoic)-Metamorphosed, subequigranular, leucocratic granite with accessory biotite, Xlg and occasional accessory garnet. Unfoliated to weakly foliated, with localized zones of mylonitic fabric. Locally cuts older mylonites. Weathers light brown to locally white. Forms the dominant rock type along the northern crest of the New York Mountains. Encloses small bodies of Xa, Xmg, and wall rock gneiss to the east, and Xfp, Xbt, and possibly Xcj to the west. Intermingled with Xgnt mainly on the east flank of the range, and commonly associated with local unmapped bodies of light gray to pink pegmatite. Miller and Wooden (1993) report U-Pb ages ranging from 1,672 to 1,695 Ma. Leucocratic granite with garnet clots (Paleoproterozoic)—Metamorphosed, subequigranular leucocratic granite. Lithologically similar to Xlg, except contains prominent clots of garnet up to 2 cm; exhibits a glomeroporphyroblastic texture in outcrop. Weathers to light brown. Exposed primarily on the eastern flank of the New York Mountains in the Crescent Peak Quadrangle, where it is complexly intermingled with Xlg and exhibits similar relationships with adjacent units as with Xlg. Porphyritic Leucocratic granite (Paleoproterozoic)—Metamorphosed, leucocratic granite with up to 2 cm-sized feldspar porphyroclasts. Feldspar porphyroclasts are commonly rounded during metamorphism and tectonic deformation, and variably overprinted by mylonitic fabric. Unit weathers light reddish brown to lavender gray. Exposed primarily in the northwestern portion of the Crescent Peak Quadrangle near the Nevada border. Locally brecciated and faulted in the northwestern-most exposures, where fluorite was locally observed in a prospect pit. Brecciated Porphyritic Leucocratic granite (Paleoproterozoic) — Brecciated Sub-type of Xlgp, typically weathers to reddish-brown with iron oxide alteration. Xlgf Leucocratic fine grained to porphyritic rock (Paleoproterozoic)—Local facies of Xlg consisting of resistant, light greenish-gray to yellowish-gray weathering, fine-grained silicic rock with local feldspar phenocrysts less than 1 cm in maximum dimension; feldspar grains range from euhedral to anhedral and are only weakly deformed. Texture and lithology suggest this rock may have originally had a silicic volcanic protolith. Underlies several fault-bounded knobs on the west flank of the New York Mountains in the Crescent Peak Quadrangle. Feldspar porphyry in fine grained matrix (Paleoproterozoic)-Medium- to dark-gray or greenish-gray rock with 1–2 cm feldspar porphyroclasts in a fine-grained matrix. Feldspar porphyroclasts are commonly rounded during metamorphism and tectonic deformation. This unit commonly forms narrow bands that are overprinted by mylonitic fabric, which may have reduced the grain size of the matrix. Mapped mainly on the west flank of the New York Mountains, where mapped bodies typically include inter-layers of Xlg and Xa; found in proximity to possible Xbt and may be genetically related. **Big Tiger Wash Sequence (Paleoproterozoic)** Granodiorite of Big Tiger Wash (Paleoproterozoic)-Metamorphosed, porphyritic biotite granodiorite. Potassium feldspar phenocrysts, up to 3 cm, commonly form euhedral laths, with a laterally variable, crude preferred alignment/foliation. The alignment may represent inherited magmatic foliation or later alignment during metamorphism. Weathers to brown color. Named after exposures in the type locality at Big Tiger Wash in Nevada Miller and Wooden, 1993). Locally mapped as small bodies in the Crescent Peak Quadrangle on inselbergs ne the Nevada border, and on the west flank of the New York Mountains. A U-Pb age of approximately 1,675 Ma was reported by Wooden and Miller (1990). Xgd Granodiorite (Paleoproterozoic)—Dark brown, subequigranular, medium-to coarse-grained biotite granodiorite. Grades to granodiorite of Big Tiger Wash with increasing content of potassium feldspar phenocrysts. Locally foliated (Miller and Wooden, 1993). Restricted to a small locality near the eastern portion of the Crescent Peak Quadrangle, and nearby Xd. Xd Diorite (Paleoproterozoic)—Dark-brown to black, medium-grained hornblende diorite; spatially associated with granodiorite of Big Tiger Wash (Miller and Wooden, 1993). Restricted to a small locality in the eastern portion of the Crescent Peak Quadrangle and associated with Xgd. Mesotype granite (Paleoproterozoic)—Gray, salt-and-pepper appearance, subequigranular, mesotype biotite granite with common fine-grained accessory garnets. Exposed as numerous small, isolated pods and larger, mappable bodies near the Nevada border within Xlg and Xgnt on the east flank of New York Mountains in the Crescent Peak Quadrangle. Miller and Wooden (1993) note close association of Xmg with Xlg. Outcrops commonly display a swirled pattern of internal banding, and outcrops are spatially associated with local migmatite and gneiss. Localized, coarse recrystallization of feldspars occurs along compositional banding. In some areas this unit is cut by pegmatitic dikes and small, garnet clot bearing apophyses. Miller and Wooden (1994) report age of crystallization from U-Pb ages of zircons at about 1685 Ma. Metamorphic rocks of Willow Wash (Paleoproterozoic) Xbg Biotite-garnet gneiss (Paleoproterozoic)—Gray to brown, well-foliated, interspersed biotite garnet gneiss, quartzofeldspathic gneiss, biotite-sillimanite gneiss, and migmatite. Mineral composition ranges from a quartzo-feldspathic gneiss with 95% quartz plus feldspar and 5% biotite; to the biotite-sillimanite gneiss with 70% quartz plus feldspar, 20% biotite, 5% sillimanite, 3% garnet, and 2% opaque minerals. Other accessory minerals throughout the unit include hornblende and apatite. The unit also includes 1- to 2-m-wide, foliation-parallel, medium grained phaneritic to porphyritic, quartzo-feldspathic dikes or leucosomes, with and without garnet. This unit forms the dominant rock type in the Willow Wash and Ivanpah areas. The unit becomes increasingly mylonitized to the east near the contact with Xgs. The mylonitization produced augen shaped, feldspar porphyroclasts and an augen-gneissic texture in this area. Miller and Wooden (1994) report U-Pb and Pb-Pb ages of zircons ranging from 1.7 Ga to 1.9 Ga. Xgs Granitoid gneiss and schist (Paleoproterozoic)—Interspersed poorly foliated quartzo-feldspathic gneiss, schistose biotite gneiss, and migmatite. The quartzo-feldspathic gneiss is gray, poorly foliated, and comprised primarily of quartz and feldspar, with 5 to 15% biotite and occasional garnets. The limited mafic minerals and lack of mineral segregation contribute to the poor foliation. The schistose biotite gneiss contains quartz, feldspar, and greater amounts of biotite than the quartzo-feldspathic gneiss with equally poor mineral segregation and foliation. Both rock types have migmatitic variations that display greater mineral segregation, but also lack well developed foliation. Xtg Tonalitic gneiss (Paleoproterozoic)—Migmatitic, equigranular biotite and hornblende-biotite-rich tonalitic gneiss, dominantly comprised of subhedral plagioclase feldspar and quartz. Local strong schistosity displayed in biotite rich intervals. Includes quartzo-feldspathic gneiss and amphibolite enclaves. Amphibolite (Paleoproterozoic)—Black to salt-and-pepper appearance, massive to foliated amphibolite. The amphibolite ranges from fine- to coarse-grained, and is comprised of plagioclase, hornblende, quartz, +/clinopyroxene, +/- biotite, and rare garnets. Xa includes granulite facies mafic rocks, containing orthopyroxene (Miller and Wooden, 1993). Interlayered in all metamorphic units, mapped where large bodies are present.



#### MAP SYMBOLS

— Contact between map units – Solid where accurately located; long dash where approximately located. Fault – Solid where accurately located; long dash where approximately located; short dash where inferred; dotted where concealed; queried where identity or existence is uncertain. <u>•</u>...? Normal Fault – Balls on upper plate, solid where accurately located; long dash where approximately located; short dash where inferred; dotted where concealed; queried where identity or existence is uncertain.

Thrust Fault – Barbs on upper plate; solid where accurately located; long dash where approximately located; dotted where concealed; queried where identify or existence is uncertain  $4 - \frac{1}{2} -$ 

**◄**—<u>+</u>—<u>+</u><u>+</u><u>-</u><u>+</u><u>+</u>?····· Antiform – Solid where accurately located; long dash where approxiately located; short dash where inferred; dotted where concealed; queried where identity or existence is uncertain.

 $\sim$ 

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IRSL geochronology point (one sample)

Strike and dip of geologic structure; number indicates dip angle in degrees.

- 50 Foliation
- <sup>10</sup> Foliation, igneous
- Gneissic foliation
- Mylonitic foliation



dotted where concealed; queried where identity or existence is uncertain. Dike, undifferentiated

Diabase dike  $\sim$   $\sim$   $\sim$   $\sim$   $\sim$   $\sim$ Mylonite zon

PRELIMINARY GEOLOGIC MAP OF THE CRESCENT PEAK 7.5' QUADRANGLE, NEW YORK MOUNTAINS, SAN BERNARDINO COUNTY, CALIFORNIA PRELIMINARY GEOLOGIC MAP 22-02 California geologic maps are available at https://www.conservation.ca.gov/cgs

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California Geological Survey



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#### **AERIAL PHOTO/DIGITAL IMAGERY**

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1. Miller, D.M., and Wooden., J.L., 1993\* \*Data source covers entire quadrangle



4 8 Miles

Currently mapped quad Other New York Mountains quads Page 1 of 1

#### AUTHORSHIP DOCUMENTATION AND PRODUCT LIMITATIONS

**PUBLICATION TITLE:** Preliminary Geologic Map of the Crescent Peak 7.5' Quadrangle, New York Mountains, San Bernardino County, California: California Geological Survey Preliminary Geologic Map 22-02

**LIMITATIONS:** This map is considered preliminary, and the California Department of Conservation makes no warranties as to the suitability of this product for any given purpose. This map should not be considered as an authoritative or comprehensive source for landslide and seismic hazard data. For landslide data, please visit the California Geological Survey Landslides web page at:

https://www.conservation.ca.gov/cgs/landslides. For seismic hazard data and Zones of Required Investigation, please visit the California Geological Survey Seismic Hazards Program web page at: https://www.conservation.ca.gov/cgs/sh/program.



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